

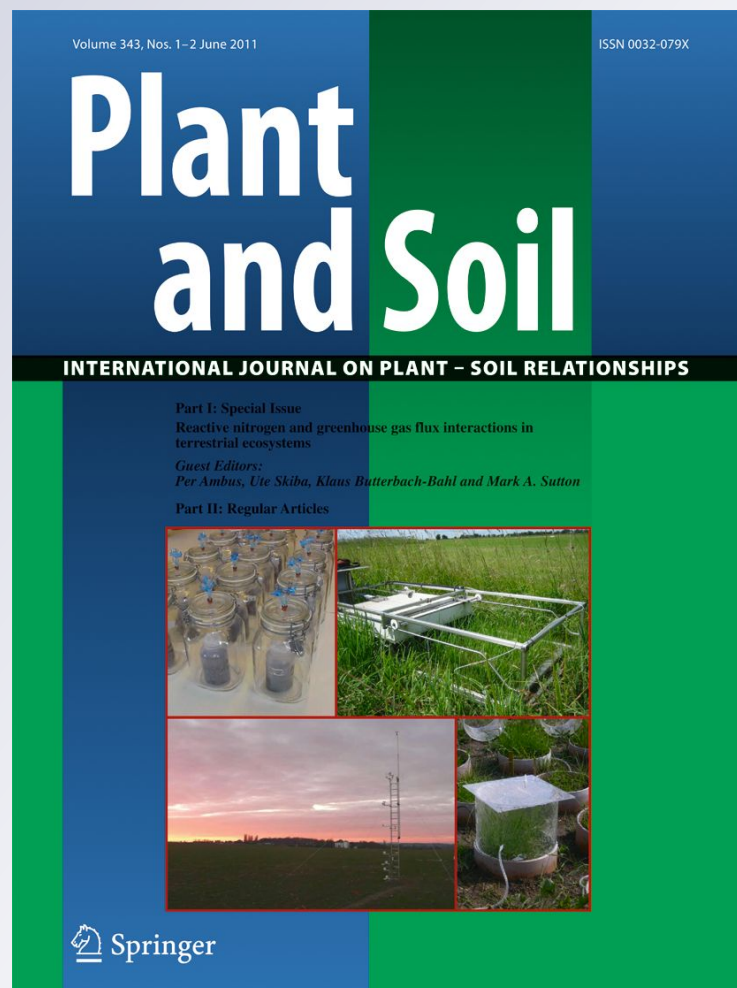
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Response of soil organic matter pools to elevated CO₂ and warming in a semi-arid grassland

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Abstract Warming and elevated atmospheric CO₂ (eCO₂) can elicit contrasting responses on different SOM pools, thus to understand the effects of combined factors it is necessary to evaluate individual pools. Over two years, we assessed responses to eCO₂ and warming of SOM pools, their susceptibility to decomposition, and whether these responses were mediated by plant inputs in a semi-arid grassland at the PHACE (Prairie Heating and CO₂ Enrichment) experiment. We used long-term soil incubations and assessed relationships between plant inputs and the responses of the labile and

resistant pools. We found strong and contrasting effects of eCO₂ and warming on the labile C pool. In 2008 labile C was increased by eCO₂ and was positively related to plant biomass. In contrast, in 2007 eCO₂ and warming had interactive effects on the labile C, and the pool size was not related to plant biomass. Effects of warming and eCO₂ in this year were consistent with-treatment effects on soil moisture and temperature and their effects on labile C decomposition. The decomposition rate of the resistant C was positively related to indicators of plant C inputs. Our approach demonstrated that SOM pools in this grassland can have early and contrasting responses to climate change factors. The labile C pool in the mixed-grass prairie was highly responsive to eCO₂ and warming but the factors behind such responses were highly dynamic across years. Results suggest that in this grassland the resistant C pool could be negatively affected by increases in plant-production driven available soil C.

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Introduction

There is a high degree of uncertainty on the combined effects of climate warming and atmospheric CO₂ increase on soil organic matter (SOM) dynamics

(Pendall et al. 2004; Shen et al. 2009). Through their effects on plant productivity, warming and eCO₂ can impact SOM dynamics, both by directly determining inputs and/or by affecting rates of losses. Changes in plant productivity and biomass turnover can alter SOM pools by affecting the rates at which below-ground production is incorporated into SOM via root death, slough off and exudation (Crow et al. 2009; Belay-Tedla et al. 2009). Increasing CO₂ generally increases above and belowground biomass (Morgan et al. 2004a; Jastrow et al. 2000; de Graaff et al. 2006) as well as root growth rate and turnover (Allard et al. 2005; Norby et al. 2004; Milchunas et al. 2005). Various meta-analyses have found that eCO₂ generally increases total soil C pools (Luo et al. 2006; Hungate et al. 2009; Jastrow et al. 2005). Warmer temperatures can also increase or limit above and belowground production, biomass and root turnover (Belay-Tedla et al. 2009; Rustad et al. 2001; Wan et al. 2004, 2005). The flux of recent photosynthate associated with increased productivity and turnover can be a major source of readily available C (Hutsch et al. 2002; Weintraub et al. 2007). For example, recent photosynthate has been found to contribute about 40% of the dissolved organic C (DOC) pool (Högberg and Högberg 2002; Giesler et al. 2007). Increases in the availability of plant-derived labile C due to CO₂ elevation and/or warming may stimulate SOM decomposition and reduce the stock of soil C, an effect known as priming (Dijkstra and Cheng 2007a; Fontaine et al. 2007; Kuzyakov 2002). Stimulation of SOM decomposition has been found to be correlated to root biomass (Fu et al. 2002) and leaf biomass (Dijkstra et al. 2006) suggesting that the extent of the stimulation could be controlled by the magnitude of the effect of CO₂ and warming on plant productivity.

Warming and eCO₂ can have contrasting effects on different SOM pools. Cardon et al. (2001) showed that eCO₂ retarded the decay of resistant, slow-turnover SOM while it increased the decomposition of labile, fast-turnover SOM. Cheng et al. (2007) observed that under eCO₂ more plant derived C was stored in the resistant pool and that its decay rate decreased. In contrast, Gill et al. (2002) found plant growth at increased CO₂ led to losses in resistant SOM, as measured by mineral-associated soil C. Other studies have observed that eCO₂ depletes both resistant SOM as well as fractions dominated by fresh

plant derived C (de Graaff et al. 2009; Del Galdo et al. 2006). Because the decomposition of labile SOM fractions is theoretically less sensitive to temperature than that of more resistant fractions (Davidson and Janssens 2006) it is likely that the effects of warming will also depend on the SOM pool examined (e.g. Belay-Tedla et al. 2009). Further, the regulation of SOM dynamics by soil moisture might not be equivalent for different pools (Garten et al. 2009). To understand the mechanisms underlying the combined effects of multiple global change factors on long-term soil C storage, effects on individual C pools and their kinetics should be evaluated.

The direction and magnitude of combined effects of eCO₂ and warming are likely to be dependent on ecosystem type (Norby and Luo 2004; Luo et al. 2008). In moisture limited systems such as semi-arid grasslands, the impacts of CO₂ and temperature on soil C are likely to be mediated by their effects on soil water availability, which can regulate both decomposition and plant productivity and thus plant inputs. Warming may induce soil drying (Rustad et al. 2001), potentially limiting plant productivity and SOM decomposition (Liu et al. 2009). However, increases in soil moisture associated with enhanced plant water use efficiency under eCO₂ (Morgan et al. 2004b) could counteract this effect. For semi-arid grasslands, models predict strong responses of plant production and soil C storage to CO₂ elevation and warming as well as interactive effects of the combined factors (Parton et al. 2007; Pepper et al. 2005; Luo et al. 2008), but empirical assessments are lacking. Our field experiment (PHACE: **P**rairie **H**eating and **C**O₂**E**nrichment), in a semi-arid, native northern mixed grass prairie in Wyoming, USA, was modeled in advance by Parton et al. (2007). Modeling results spanning ten years after the beginning of treatments predict a high dependency of soil C flux on yearly variations in soil moisture availability. A gradual decrease in soil organic C is predicted for plots exposed to eCO₂ (with or without warming) due to faster decomposition associated with higher soil moisture under eCO₂, which would surpass the effects of increased plant C input.

Here, we present an empirical evaluation of the impacts of eCO₂ and warming on SOM at the PHACE experiment. We investigated two main questions: 1) how do eCO₂, warming and their combination affect SOM pools and their susceptibility

to decomposition? and 2) are the responses of SOM pools and their susceptibility to decomposition associated with plant inputs? To address these questions, we conducted long-term laboratory incubations because they provide a more sensitive indicator of SOM responses to climate change than bulk soil measurements (Pendall and King 2007; Langley et al. 2009). Incubations under optimal conditions allow assessments of SOM pools and kinetics while removing the environmental constraints and confounding effects at play in a field situation (Davidson and Janssens 2006). The rate of microbial respiration during laboratory incubations and the kinetic parameters that can be derived from them can be used to assess changes in the susceptibility of C pools to decomposition due to previous field exposure to experimental treatments (Taneva and Gonzalez-Meler 2008; Collins et al. 2000; Carney et al. 2007). In order to examine the role of plant inputs on SOM pool sizes and their susceptibility to decomposition, we assessed whether there were relationships between indicators of plant inputs (aboveground biomass, fine root biomass and DOC measured at the time of soil collection) and the responses of the labile and resistant pools.

Materials and methods

Study site

The Prairie Heating and CO₂ Enrichment (PHACE) experiment at the USDA-ARS High Plains Grassland Research Station (HPGRS) is located about 15 km west of Cheyenne, WY, USA (41° 11' N, 104° 54' W; 1930 m. a.s.l.). The ecosystem is a northern mixed grass prairie with a plant community dominated by the C4 grass *Bouteloua gracilis* (H.B.K) Lag. and two C3 grasses, *Hesperostipa comata* Trinand Ruprand and *Pascopyrum smithii* (Rydb.). About 20% of the vegetation is composed of sedges and forbs. Annual precipitation is 384 mm, with ca. 60% falling during the growing season, which extends from late March through September. Mean air temperatures are 17.5°C in summer and -2.5°C in winter, with average maximum July temperature of 27°C (meteorological data collected on-site and compiled by HPGRS). Soils at the experimental site are Mollisols (fine-loamy, mesic Aridic Argiustoll, mixed Ascalon and Altvan series), with a pH of 7.0. Average

organic soil C content is 1.9% (S.D.=0.27) at 0–5 cm and 1.5% (S.D.=0.25) at 5–15 cm. Average carbonate content by weight is 0.0045 % at 0–5 cm (S.D.=0.062) and 0.2056 % (S.D.=0.61) at 5–15 cm. The experiment uses Free Air CO₂ Enrichment (FACE) technology (Miglietta et al. 2001) installed in 3.4-m diameter rings to produce an atmosphere with a CO₂ concentration of 600±40 μmol mol⁻¹ in the eCO₂ treatments during the growing season. Infrared heaters (Kimball et al. 2008) attached to a frame 1.5 m above the ground increase air temperature by 1.5°C during daytime and 3°C at night all year round. The plots are thus exposed to four combinations of CO₂ and temperature (ambient CO₂ and ambient temperature: ct; ambient CO₂ and elevated temperature: cT; eCO₂ and ambient temperature: Ct; eCO₂ and elevated temperature: CT) with five replications. CO₂ enrichment began at the start of the growing season (early April) in 2006 and heating began at the start of the growing season in 2007.

Field measurements and sample processing

In each experimental plot, soil volumetric water content is monitored hourly at 10 and 20 cm (EnviroSMART sensors, Sentek Sensor Technologies, Stepney, Australia). Soil temperature is also measured hourly at 3 and 10 cm using thermocouples. Daily means of soil water content and temperature were averaged across the six-week periods prior to soil sampling (Shim et al. 2009). Standing fine root biomass and aboveground plant biomass were assessed at the peak of the growing season (late July, close to the time of soil collection) in 2007 and 2008. Fine roots (<1 mm) were removed from soil samples (see below). Aboveground biomass was obtained by clipping an area of 0.75 m² in each plot.

In June 2007 and July 2008 three soil cores (15-cm deep; 3-cm diameter) were collected from each plot, divided into 0–5 and 5–15 cm depths, consistent with pretreatment protocols, and composited into one sample per depth. Fresh soils were sieved to 2 mm and further inspected to remove visible litter and root material. To obtain a measure of DOC, a 15-g subsample of fresh soil was extracted by shaking in 30 ml of 0.05 M K₂SO₄ for 1 hour. The filtered extract was analyzed for total C (Shimadzu TOC-V_{CPN}, Shimadzu Scientific Instruments, Wood Dale, IL). Bulk soil organic C content was determined for a subsample

of air-dried soil which was ground to powder and analyzed with a Costech EA 1108 Element Analyzer. For incubations we used soil from 0 cm to 15 cm in depth, which contains very low and mostly undetectable levels of carbonates (see [Study site](#)). However, prior to analyses, and for consistency, all soil samples and extracts from our experiment are routinely treated to remove carbonates as deeper soils have higher contents that require removal. Extracts were treated with 1 M H_3PO_4 (1 μl per 10 ml of extract) and soil samples were shaken with 1 M H_3PO_4 (Sherrod et al. 2002).

Incubations

On the day after collection, subsamples (15–20 g) of root/litter-free 2-mm sieved soil were placed in polystyrene beakers which had small holes punched into the bases and pre-combusted glass fiber filters in the bottom for drainage (Townsend et al. 1997). Sub-samples of soils collected from three and five experimental replicates were incubated in 2007 and 2008 respectively. Deionized water was added to each beaker to reach 60% of the soil's field capacity. Each beaker was placed into a 500-ml canning jar with a lid modified to hold a 1.5-cm, blue butyl rubber stopper. Soils were incubated at 25°C in a controlled environment incubator. Headspace samples (30 ml) were collected at 1, 2, 5, 8, 9, 12, 22, 33, 43, 79 and 107 days of incubation in 2007 and at 1, 2, 5, 9, 13, 22, 30, 45, 78 and 114 days in 2008. Prior to headspace sampling jars were opened to allow equilibration with ambient air and then closed and flushed with CO_2 -free air for 2 min to bring the headspace CO_2 concentration to zero. Jars were then immediately placed in the incubator for periods ranging from 24 h during the early days of incubation to 50 hours at the final sampling date, to allow approximately 1000 $\mu\text{mol mol}^{-1}$ CO_2 to accumulate. CO_2 production rate per hour was calculated based on the length of time after flushing. Four empty jars were used as controls. Headspace samples were analyzed for CO_2 concentration with an infrared gas analyzer (Li-Cor 820, LICOR Inc. Lincoln, NE) calibrated with four standard gases over the concentration range 377–4500 $\mu\text{mol mol}^{-1}$.

To estimate C kinetic parameters, measured CO_2 production rates over time were fitted to a two pool decay model in which a pool of labile C decays exponentially and a pool of resistant C decays at a

constant rate (Dijkstra and Cheng 2007a; Wedin and Pastor 1993):

$$R_t = R_i * e^{-k_i * t} + R_r$$

Where R_t = respiration rate at time t ($\mu\text{g C g soil}^{-1} \text{ day}^{-1}$); R_i = initial rate of decomposition of the labile pool (at the start of the incubation; $\text{mg C g soil}^{-1} \text{ day}^{-1}$); k_i = decay rate of C in the labile pool (day^{-1}); R_r = decomposition rate of C in resistant pool ($\text{mg C g soil}^{-1} \text{ day}^{-1}$). The size of the labile pool (C_l) was defined as R_i/k_i . Curve fitting was done using the single, 3 parameter exponential decay equation in Sigma Plot 10.0. To account for pre-experimental soil C conditions, R_i , C_l and R_r were standardized by dividing the estimated values by total soil C content, measured for each plot and at each depth. The labile pool is reported as $\text{mg C g soil C}^{-1}$ and the decomposition rates of the labile and resistant pools as $\text{mg C g soil C}^{-1} \text{ day}^{-1}$.

Statistical analyses

To determine the effects of depth, CO_2 and warming treatments on field soil moisture, temperature, and estimated parameters (size of the labile C pool, the rates of decomposition of the labile and resistant pool) over the course of incubation we used a split plot, three-factor ANOVA with CO_2 , warming and depth as fixed effects and replicate and plot \times depth as random effects. Each one of the incubations was analyzed as an independent experiment. We preferred this approach to a single model analysis as there was a small (6%) difference in the duration of the incubation and slightly more rapid drying of the 2007 soils in the last few weeks of incubation. These differences could have influenced the magnitude of some of the estimated parameters across all treatments, but were not expected to have influenced treatment effects. Consequently, our analyses and interpretations focus on the treatment effects observed in each year and the variables driving them rather than on the comparison of absolute magnitudes of the pools and rates across years. Above-ground plant biomass, root biomass, DOC and average daily soil water content means for the six weeks prior to soil sampling were linearly regressed against the estimated pool sizes and kinetic parameters. Significant effects and relationships are reported at $p < 0.05$ unless otherwise stated. All

statistical analyses were performed with JMP (version 8.0; SAS Institute, Cary, N.C. USA).

Results

Effects of field treatments on soil temperature and moisture

Average soil volumetric water content ranged from 8 to 13% at 0–10 cm depth, and 12 to 18% at 10–20 cm depth (Table 1). Over the six weeks previous to mid-season sampling in 2007 and 2008, soil water content was significantly greater (by 2.1% water content on average; Table 1) in eCO₂ than ambient CO₂ plots ($p=0.03$ for 2007 and $p=0.002$ for 2008). Warming significantly decreased water content prior to all sampling dates by 2.0% on average ($p=0.04$ for 2007 and $p=0.006$ for 2008; Table 1). No interactions between treatment and depth were observed for soil moisture.

eCO₂ significantly decreased soil temperature by 0.2°C in the period before sampling in 2007 and 0.3°C in 2008 ($p=0.07$, $p=0.0001$ respectively). Warming significantly increased soil temperature during all periods (Table 1). The impact of warming depended on soil depth (significant interaction of warming and soil depth, $p=0.0004$ for 2007 and $p=0.0003$ for 2008), as the increase in soil temperature was greater at 3 cm than at 10 cm (2.8°C vs 2.1°C on average in 2007 and 2.1°C vs 1.8°C in 2008).

Plant biomass and DOC

eCO₂ increased fine root biomass in 2007 and 2008 and aboveground biomass in 2008 (Table 2). There

was no main effect of warming, but the combination of eCO₂ and warming tended to enhance fine root biomass in 2008 (marginally significant CO₂ x warming interaction; Table 2). eCO₂ had no detectable effect on DOC while warming tended to increase DOC concentration in 2008 (Table 2). Aboveground plant biomass and fine root biomass were positively correlated (Fig. 1). The concentration of DOC in soil was positively correlated to both aboveground plant biomass and root biomass (Fig. 2).

Estimated pool sizes and decomposition rates

The proportion of labile C and its initial decomposition rates showed strong interactive effects of CO₂ and warming in the 0–5 cm layer in soils sampled in 2007. Warming increased the proportion of total C that was labile in soils from ambient CO₂ plots but decreased it in eCO₂ soils (Fig. 3a). In 2007 the decomposition rate of the labile C was lower with warming in ambient CO₂ soils, while it increased under warming in eCO₂ soils (Fig. 3c). In 2008, the labile proportion was greater in eCO₂ soils and this effect was consistent across depths (Fig. 3b). The resistant pool decomposition rate was not significantly affected by the CO₂ and warming treatments or by depth in either year (Fig. 3e, f).

The proportion of labile C was positively related to aboveground biomass at both 0–5 cm and 5–15 cm depths and to fine root biomass at 0–5 cm in 2008 (Fig. 4). In contrast, the labile C in soils from 2007 was not related to plant biomass or DOC (regression plots not shown). The initial rate of decomposition of the labile pool was positively related to soil water content at both depths ($p=0.04$, $r^2=0.34$ at 0–

Table 1 Average daily soil water content (volumetric) and temperature at two soil depths over the six week period prior to soil collections in 2007 and 2008

Values are means with standard error in parentheses. *ct* ambient CO₂ and ambient temperature; *cT* ambient CO₂ and elevated temperature; *Ct* elevated CO₂ and ambient temperature; *CT* elevated CO₂ and elevated temperature

	2007		2008	
	0–10 cm	10–20 cm	0–10 cm	10–20 cm
Soil water content (%)				
<i>ct</i>	10.6 (0.59)	15.6 (1.00)	8.9 (0.25)	13.8 (0.82)
<i>cT</i>	9.7 (0.38)	12.7 (0.84)	7.9 (0.22)	11.5 (0.72)
<i>Ct</i>	13.0 (0.87)	17.7 (1.35)	11.6 (0.63)	16.2 (1.08)
<i>CT</i>	11.5 (0.67)	15.1 (1.30)	9.6 (0.57)	13.6 (1.25)
Soil temperature (°C)				
<i>ct</i>	19.0 (0.17)	17.5 (0.16)	24.9 (0.15)	23.2 (0.15)
<i>cT</i>	21.6 (0.12)	19.5 (0.06)	28.4 (0.11)	25.2 (0.22)
<i>Ct</i>	18.6 (0.10)	17.3 (0.12)	24.1 (0.11)	22.7 (0.20)
<i>CT</i>	21.5 (0.20)	19.4 (0.07)	27.6 (0.31)	25.0 (0.13)

Table 2 Above-ground biomass (g m^{-2}), fine root biomass (g m^{-2} of ash-free dry weight in top 15 cm), and dissolved organic C in soil (DOC, g m^{-2} in top 15 cm) at the time of peak aboveground biomass in 2007 and 2008

	Above-ground biomass		Fine root biomass		DOC	
	2007	2008	2007	2008	2007	2008
ct	102.2 (5.4)	87.9 (1.7)	262.6 (27.6)	344.2 (27.4)	3.2 (0.4)	4.4 (0.4)
cT	104.6 (9.8)	89.9 (5.8)	215.1 (25.6)	294.6 (28.1)	2.7 (0.3)	5.3 (0.6)
Ct	99.7 (7.7)	95.7 (8.2)	295.8 (34.6)	342.2 (31.2)	3.1 (0.5)	4.8 (0.5)
CT	132.1 (13.2)	104.8 (4.7)	302.1 (29.5)	441.7 (60.4)	3.2 (0.5)	6.5 (1.0)
ANOVA p-values						
CO ₂	0.21	0.005	0.07	0.08	0.64	0.24
Warming	0.09	0.14	0.51	0.53	0.64	0.06
CO ₂ x Warm.	0.13	0.35	0.39	0.08	0.48	0.56

Values are means with standard error in parentheses. *ct* ambient CO₂ and ambient temperature; *cT* ambient CO₂ and elevated temperature; *Ct* elevated CO₂ and ambient temperature; *CT* elevated CO₂ and elevated temperature

5 cm and $p=0.03$, $r^2=0.39$ at 5–15 cm) in soils from 2007, and at 5–15 cm in 2008 ($p=0.02$, $r^2=0.27$) (regression plots not shown). Significant relationships between the resistant pool decomposition rate and the indicators of plant input were observed in both years: the resistant pool decomposition rate was positively related to aboveground plant biomass, root biomass and DOC in 2007 and with DOC in 2008 (Fig. 5).

Discussion

We evaluated the responses to eCO₂ and warming of SOM pools and their susceptibility to decomposition in a semi-arid grassland and whether these responses were mediated by plant inputs. We used aboveground

biomass (in this system, equivalent to aboveground net primary productivity) and fine root biomass as measures of plant inputs to soil. Our results demonstrated that positive aboveground productivity responses to eCO₂ and warming treatments translated into greater belowground biomass, which in turn was associated with an enlarged pool of soil DOC (Figs. 1 and 2). An important source of DOC is root exudation (Högberg and Högberg 2002; Giesler et al. 2007), which is, in turn, proportional to root biomass (Jones et al. 2009) and fixed C (Pinton et al. 2001). Together, these results indicate that aboveground biomass, fine root biomass and DOC could be used as indicators of plant C inputs to soil.

The marked treatment effects on the labile C observed in both years demonstrated the sensitive nature of this pool to CO₂ and warming. Contrasting effects across years indicate the dynamic nature of the factors involved in the regulation of this pool. Our results suggest that plant input was more important in determining the size of the labile pool in 2008 while treatment effects on soil moisture and temperature and their impacts on decomposition dominated the effects in 2007. The greater importance of plant inputs in 2008 is supported by positive relationships between the labile C pool size and plant biomass (Fig. 4). A greater role of plant inputs in controlling the labile C pool in 2008 than in 2007 is also consistent with the positive effect of eCO₂ on aboveground biomass observed in 2008 (Table 2). A significant increase in aboveground biomass is expected to increase the input of recent photosynthate to soil which apparently enlarged the

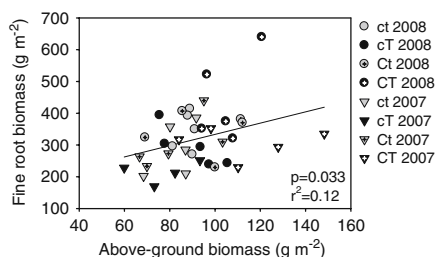


Fig. 1 Linear relationship between above-ground plant biomass and fine root biomass (0–15 cm depth) at the time of peak biomass (late July) in 2007 and 2008. *ct*: ambient CO₂ and ambient temperature; *cT*: ambient CO₂ and elevated temperature; *Ct*: elevated CO₂ and ambient temperature; *CT*: elevated CO₂ and elevated temperature

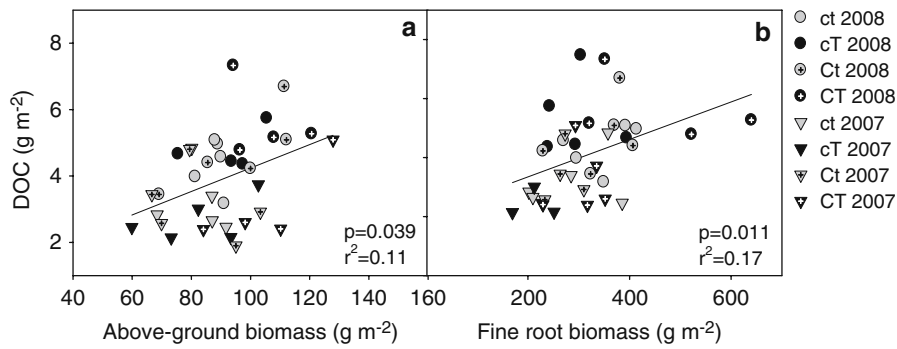


Fig. 2 Linear relationship between field measurements of dissolved organic C in soil (0–15 cm; DOC) and above-ground plant biomass **a** and fine root biomass (0–15 cm) **b** at the time of peak biomass in 2007 and 2008. **ct**: ambient CO₂

and ambient temperature; **cT**: ambient CO₂ and elevated temperature; **Ct**: elevated CO₂ and ambient temperature; **CT**: elevated CO₂ and elevated temperature

Fig. 3 Labile pool size (**a**, **b**), its initial decomposition rate (**c**, **d**) and the decomposition rate of the resistant pool (**e**, **f**) in soils collected in June 2007 and July 2008. Values were estimated through non-linear regression of observed respiration rates under incubation conditions (see Methods) and are presented per gram of soil C to account for pre-experimental soil C concentrations. Values are means standard error. **ct**: ambient CO₂ and ambient temperature; **cT**: ambient CO₂ and elevated temperature; **Ct**: elevated CO₂ and ambient temperature; **CT**: elevated CO₂ and elevated temperature. p values of treatment effects from split-plot 3-factor ANOVA (CO₂, warming and depth) are shown only when effects were considered significant

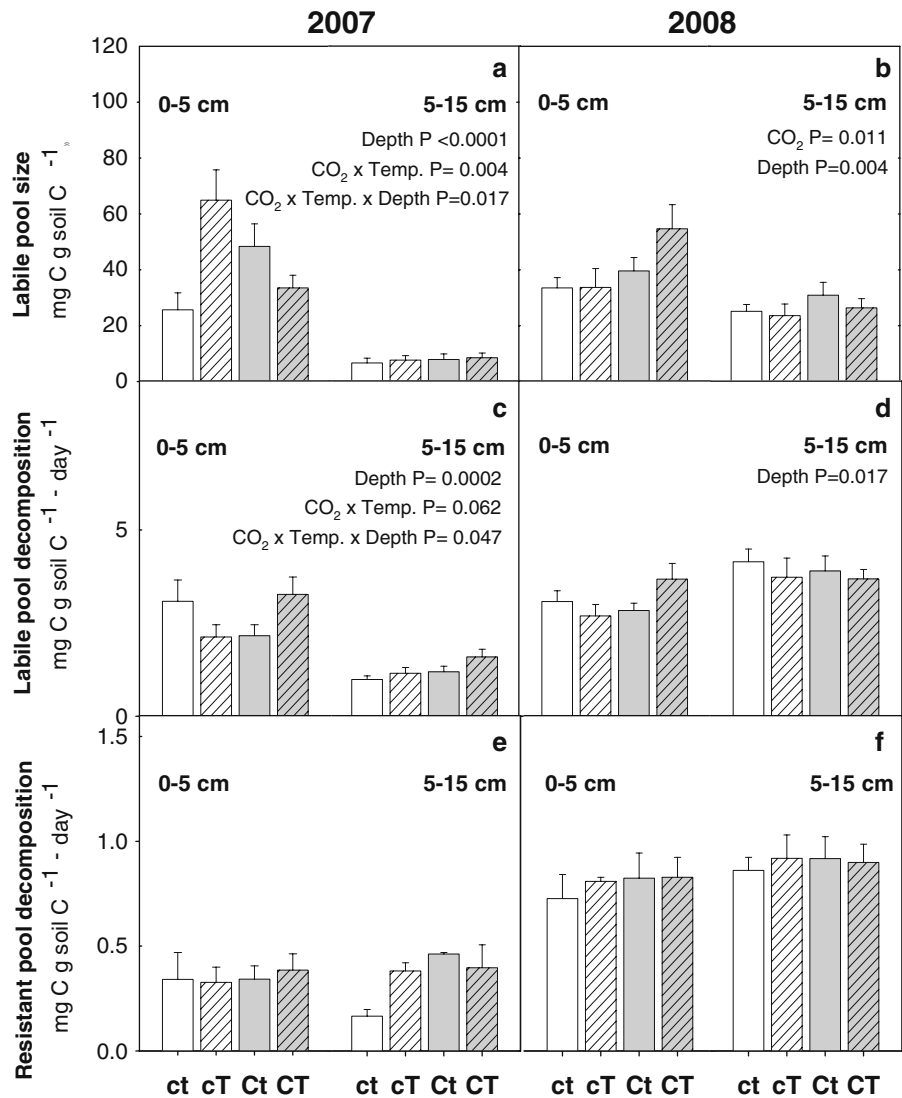
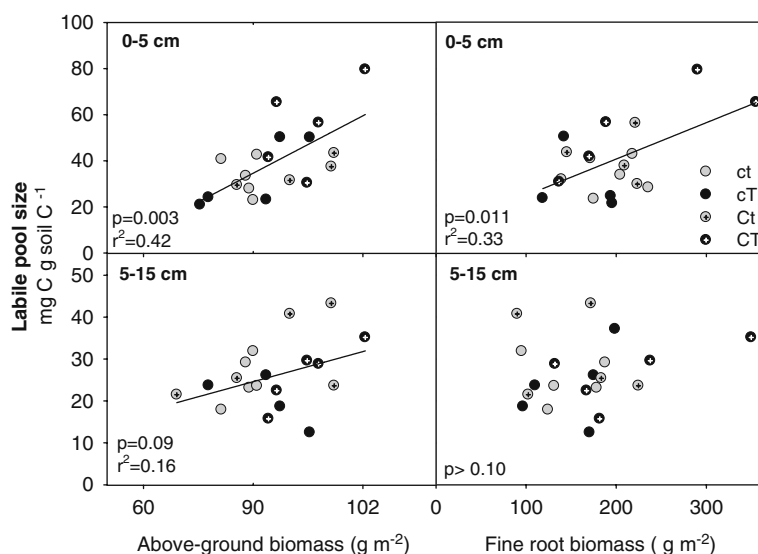


Fig. 4 Linear relationships between the estimated pool of labile C at 0–5 cm and 5–15 cm soils (as mg C g soil C⁻¹) and field measured above-ground biomass and fine root biomass for soils collected close to the time of aboveground peak biomass in 2008. **ct**: ambient CO₂ and ambient temperature; **cT**: ambient CO₂ and elevated temperature; **Ct**: elevated CO₂ and ambient temperature; **CT**: elevated CO₂ and elevated temperature



labile C pool in 2008. In contrast, in 2007, labile C was not related to plant biomass variables and there were no strong effects of treatments on plant biomass. Observed interactive effects of warming and CO₂ on the labile C pool in 2007 are, instead, consistent with the effects of eCO₂ and warming on soil moisture and temperature and with their effects on decomposition rates. The higher content of labile C in ambient CO₂ with warming (Fig. 3a; cT) in 2007 could be explained by decreased labile C decomposition by

microbes due to soil drying (Liu et al. 2009). In eCO₂ with warming (Fig. 3a; CT), the effect of drying on the labile pool was likely offset by greater soil moisture due to higher plant water use efficiency under eCO₂ (Table 1) (Morgan et al. 2004b), thus allowing enhanced decomposition of the labile C with increased temperature (Kirschbaum 2006).

When considering all treatments in each year, the decomposition rate of the resistant C pool was positively related to plant (above and belowground)

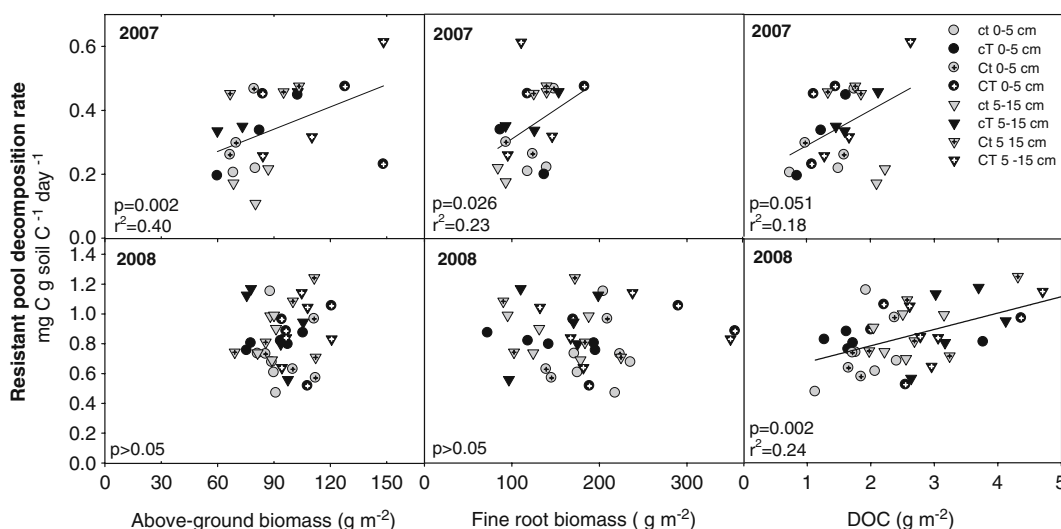


Fig. 5 Linear relationships between the estimated rate of decomposition of the resistant C pool from 0–5 and 5–15 cm soils (as mg C g soil C⁻¹ day⁻¹) and field measured aboveground biomass, root biomass and dissolved organic C

(DOC) for soils collected close to the time of aboveground peak biomass in 2008. **ct**: ambient CO₂ and ambient temperature; **cT**: ambient CO₂ and elevated temperature; **Ct**: elevated CO₂ and ambient temperature; **CT**: elevated CO₂ and elevated temperature

biomass in 2007 and to the concentration of DOC in soil in 2007 and 2008 (Fig. 5). Although in 2008 DOC was the only variable with detectable relationships with the resistant pool, both root and aboveground biomass were positively related to DOC (Fig. 2) and labile C in this year (Fig. 4). The positive correlations between plant biomass variables with DOC and labile C suggest that enhanced plant inputs increased the concentration of readily available C in soil. Together, these observations suggest that, at least under incubation conditions, the susceptibility to decomposition of the resistant pool was greater in soils that had greater concentrations of plant derived inputs and thus greater concentrations of available C. Increased available C in soil has been linked to enhanced decomposition of SOM via the priming effect (Fontaine et al. 2007; Kuzyakov 2010). Thus, the positive relationships between the decomposition rate of the resistant pool with plant biomass and DOC that we observed suggest that greater concentrations of readily available C in soil could have made the resistant C pool more susceptible to decomposition, in other words, that the resistant pool in these soils has the potential to be primed by increased labile C under field conditions.

A disadvantage of incubation conditions is that soil C pools are decoupled from the root system. Thus it is not possible to directly measure stimulation of SOM decomposition resulting from root C inputs, or 'rhizosphere priming' (Kuzyakov 2002). However, by measuring aboveground and root biomass as well as DOC at the time of soil collection and immediately setting up controlled incubations of freshly collected soils to determine labile C pools we were able to link field treatment impacts on plant inputs with C availability in soil. Stimulation/suppression of SOM decomposition is often assessed by measuring total CO₂ fluxes and not directly by assessing decomposition rates of individual pools and thus it is difficult to determine whether the extra C respired is derived from microbial turnover, labile soil C or stable OM (Blagodatskaya and Kuzyakov 2008). However, a change in the dynamics of long-lived, resistant SOM pools has the highest potential to impact long term C storage. With an incubation approach, we were able to estimate the rate of decomposition of the resistant pool, in particular, as opposed to bulk SOM, under the range of labile C availability generated by the experimental treatments.

eCO₂ has been hypothesized to cause SOM depletion via the priming effect because it can promote plant production and thus plant C input to soil (Carney et al. 2007). Studies have found evidence of accelerated decomposition of older SOM under eCO₂ (Pendall et al. 2003) as well as of pools of total and resistant SOM under eCO₂ (Langley et al. 2009; Hoosbeek et al. 2004), which have been attributed to priming. Other climate factors, however, might contribute to C depletion by altering the availability of labile C. For example, changes in moisture availability due to either warming or CO₂ could contribute to changes in labile C availability and thus impact the extent of SOM priming (Dijkstra and Cheng 2007b). Our observations under incubation conditions suggest that the stable SOM in soils from the mixed-grass prairie is sensitive to plant C inputs, which in turn were related to available C in soil. Based on the positive relationships observed between indicators of plant inputs and the decomposition rate of the resistant C at PHACE, we speculate that if increased C availability were to occur and persist over multiple growing seasons it could lead to a gradual depletion of the resistant C pool. Using the DayCent model, Parton et al. (2007) predicted that despite increased plant inputs, eCO₂ alone and in combination with warming will gradually decrease total soil organic C at PHACE, due to enhanced decomposition with greater soil moisture (a mechanism that can be accounted for by the DayCent model). Changes in the dynamics of long-lived, resistant SOM pools have the highest potential to impact long term C storage. Our results suggest an additional mechanism by which soil C, the resistant pool in particular, could be depleted in this system, highlighting the need to include the influence of labile C on the dynamics of long-lived, resistant SOM pool, in the modeling of long term responses of soil pools to climate change (Heimann and Reichstein 2008; Chapin et al. 2009).

Changes in bulk soil C pools are usually detected only after long exposures to global change treatments. By assessing individual C pools, our study demonstrated that SOM pools in the semi-arid mixed grass prairie can have early responses to eCO₂ and warming and that different pools can show contrasting effects. Our results demonstrated that the labile C pool is highly responsive to eCO₂ and warming but that the main factors driving the responses of this pool are highly dynamic. Plant inputs were more important in

controlling the size of the labile pool in the year when aboveground biomass responded to eCO₂, which highlighted the strong link between plant biomass responses to climate factors and labile C availability in soil. Our assessment of the decomposition rate of the resistant C pool suggested that the stable SOM in soils from the mixed-grass prairie is sensitive to C availability and particularly that it could be negatively affected by increases in readily available C driven by global change impacts on plant productivity. An in situ evaluation of the response of resistant SOM under field conditions is thus warranted.

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